EVALUATION OF POTENTIAL HYDROCARBON TRANSPORT AT THE UC-4 EMPLACEMENT HOLE, CENTRAL NEVADA TEST AREA

prepared by

Brad F. Lyles, Greg Pohll, Derek Sloop, Charalambos Papelis

submitted to

Nevada Operations Office
U.S. Department of Energy

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prepared by

Brad F. Lyles, Greg Pohll, Derek Sloop, Charalambos Papelis
Water Resources Center
Desert Research Institute
University and Community College System of Nevada

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Nevada Operations Office
U.S. Department of Energy
Las Vegas, Nevada

September 1998

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ABSTRACT

Emplacement hole UC-4 was drilled in 1969 at the Central Nevada Test Area and left filled with drilling mud. Surface characterization samples collected from abandoned mud pits in the area yielded elevated concentrations of total petroleum hydrocarbon, thereby raising a concern that the mud-filled emplacement hole may be leaching hydrocarbons into alluvial aquifers. This study was initiated to address this concern.

An analytical solution for flow near a wellbore was used to calculate the amount of time it would take for a contaminant to move through the mud-filled well and into the surrounding aquifer. No hydraulic data are available from the emplacement hole; therefore, ranges of hydraulic conductivity and porosity were used in 100 Monte Carlo realizations to estimate a median travel time. Laboratory experiments were performed on samples collected from the central mud pit to determine the hydrocarbon release function for the bentonite drilling mud.

The median contaminant breakthrough took about 12,000 years to travel 10 m, while the initial breakthrough took about 300 years and the final breakthrough took about 33,000 years. At a distance of about 10 m away from the emplacement hole, transport velocity is dominated by the hydraulics of the aquifer and not by the emplacement hole hydraulics. It would take an additional 45,500 years for the contaminant to travel 800 m to the U.S. Department of Energy land exclusion boundary. Travel times were primarily affected by the hydraulic conductivity and porosity of the drilling mud, then by the hydraulic conductivity, porosity and hydraulic gradient of the alluvial aquifer, followed by the hydrocarbon release function.
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<th>DESCRIPTION</th>
</tr>
</thead>
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<tr>
<td>CNTA</td>
<td>Central Nevada Test Area</td>
</tr>
<tr>
<td>DOE</td>
<td>U.S. Department of Energy</td>
</tr>
<tr>
<td>EPA</td>
<td>U.S. Environmental Protection Agency</td>
</tr>
<tr>
<td>FID</td>
<td>flame ionization detector</td>
</tr>
<tr>
<td>GC</td>
<td>gas chromatograph</td>
</tr>
<tr>
<td>HPLC</td>
<td>high pressure liquid chromatography</td>
</tr>
<tr>
<td>I.D.</td>
<td>inside diameter</td>
</tr>
<tr>
<td>IT</td>
<td>International Technology Corporation</td>
</tr>
<tr>
<td>TPH</td>
<td>total petroleum hydrocarbon</td>
</tr>
<tr>
<td>USAEC</td>
<td>U.S. Atomic Energy Commission</td>
</tr>
<tr>
<td>VOA</td>
<td>volatile organic analysis</td>
</tr>
</tbody>
</table>
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INTRODUCTION

Calculations were performed to determine the nature and extent of potential groundwater contamination at well UC-4 on the Central Nevada Test Area (CNTA). Although the hydrocarbon concentration in the mud-filled borehole of UC-4 is unknown, as a conservative estimate, the highest observed concentrations of total petroleum hydrocarbon (TPH) from abandoned mud pits in the area were used. Laboratory experiments were performed to derive a release function for TPH from semi-consolidated mud samples collected from the central mud pit, near the UC-1 site. The source of TPH in the mud pit is not clear, but may be from one or more of the following sources: 1) used motor oil, 2) oil and grease from an equipment cleaning area near the mud pit, or 3) a drilling mud additive. TPH samples collected from the UC-4 mud pits had considerably lower concentration than from the central mud pit; however, as a worst-case scenario it was assumed that the same practices used at the central mud pit were used at the UC-4 mud pits. An analytic solution for groundwater flow near a borehole was used to estimate the transport of TPH. An exponential release function was used to simulate the TPH release from the drilling mud.

UC-4 DRILLING AND CHARACTERIZATION HISTORY

Emplacement hole UC-4 was drilled in the northernmost testing area of the CNTA (Figure 1). The exclusion area around the UC-4 site encompasses 3.88 km² (Public Land Order No. 4748 of 12/2/1969). UC-4 was drilled from April to September 1969, to a total depth of 1676 m and is 3 m in diameter (see Figure 2). The upper 126.5 m of the hole was completed with 3.1-m-diameter steel casing, the remainder of the well was left uncased. Portions of bit #45 were lost in the well on April 12, at a depth of 1407.2 m. Fishing and milling operations lasted for the next 30 days (Fenix and Sisson, 1968). Although no details were recorded in the hole history regarding the addition of fluids during the fishing operation, it was not uncommon to inject large quantities of diesel fuel into the borehole during fishing operations as a lubricating agent (J.D. Donithan, Desert Research Institute, pers. comm. 1998). The borehole was filled with drilling mud upon the completion of drilling. Final logging and testing of the well were completed December 5, 1969, at which time the fluid level had dropped to 79.8 m below land surface. The hole was then abandoned by welding a 5.1-cm-thick steel plate to the top of the 3.1-m-diameter steel casing; a 4.9 m by 4.9 m by 0.6 m concrete pad was then poured on top of the steel plate (USAEC, 1974). This is the current configuration of the wellhead; there is no ready way to access the wellbore to collect new information. Therefore, all calculations are made from mud pit samples and from literature-reported values.

No lithologic data can be found for well UC-4; however, detailed lithologic descriptions are available for exploration well UCe-17, located approximately 100 m west of UC-4 (Hoover, 1968). Due to the close proximity of UCe-17, it is assumed that the lithology and formation tops would be very similar to UC-4. The upper 625 m are alluvium, composed of caliche-cemented silt to cobble sized particles of volcanic detritus. The lithology from 625 m to 1676 m (total depth) is composed of four volcanic tuff units (Orange Lichen tuff, Butte tuff, Needles tuff, and Biotite tuff). The volcanic units have varying mineralogical compositions and are from low to moderate to densely welded.
Figure 1. Location of well UC-4 at the Central Nevada Test Area in Hot Creek Valley, Nevada.
Figure 2. Generalized well schematic and lithologic log for well UC-4.
Groundwater flow near well UC-4 is dominated by downward vertical flow, with a vertical gradient of approximately 0.1 (K. Pohlmann, Desert Research Institute, per. comm., 1998). However, the horizontal gradient is of greatest concern for this analysis of hydrocarbon transport. No wells are available in the immediate vicinity of UC-4 from which water level measurements can be made. Therefore, estimates of groundwater flow direction and gradient were made from previously measured water levels between well UCe-17 and current levels at HTH-1 near UC-1, approximately 5.7 km south of the UC-4 site. The horizontal gradient between UC-4 and HTH-1 was 0.02.

Transport of TPH was presumed to be most likely at the water table in the alluvial aquifer. Values of hydraulic conductivity and porosity in the alluvial aquifer were available from UCe and HTH wells in the area, though no hydraulic conductivity values were available from well UC-4, and only values from the volcanics were available from UCe-17. The hydraulic conductivity of the alluvial aquifer has a geometric mean and log10 standard deviation of 0.87 m/yr and 0.091, respectively. The arithmetic mean and standard deviation of the porosity is 19.3 percent and 6.31 percent, respectively (K. Pohlmann, Desert Research Institute, per. comm., 1998). For the computer model calculations, the hydraulic conductivity was assumed to be log-normally distributed and the porosity was assumed to be normally distributed.

No documentation is available on the mud used to fill well UC-4. However, the mud engineer remembered that the mud was composed of bentonite and chrome lignosulfonate additives (McLauren, Summit Minerals, Denver, CO (303) 572-3011, pers. comm., 1998). Drilling mud is specifically designed to stabilize the formation, hold the hole open and carry rock cuttings during drilling operations, and by its nature is not viewed as a porous medium. Therefore, no data are available on the hydraulic conductivity or porosity of drilling mud, nor are there data on the long-term effects of the drilling mud left in wells. However, it is presumed that the mud left in UC-4 has consolidated to a clay over the past 29 years since the well was completed. Therefore, rather than performing laboratory tests on drilling mud samples, literature-reported values for clays were used to estimate the hydraulic properties of the mud-filled well. The geometric mean hydraulic conductivity of the mud was chosen to be 0.002 m/yr with a log10 standard deviation of 0.0317 m/yr and the arithmetic mean porosity of the mud was 55 percent with a standard deviation of 58.5 percent (Freeze and Cherry, 1979). For the computer model calculations the distributions were assumed to be the same as for the alluvium; the hydraulic conductivity was assumed to be log-normally distributed and the porosity was assumed to be normally distributed.

**ANALYTICAL SOLUTION FOR FLOW NEAR A WELL OF LOW HYDRAULIC CONDUCTIVITY**

Groundwater flow near a well is distorted due to the change in hydraulic conductivity from the aquifer to the borehole. In most cases the hydraulic conductivity of the well is greater than that of the surrounding medium, thereby deforming the flow lines toward the well. In the case of a well filled with mud, the flow lines of equal potential are distorted away from the well (Figure 3). Wheatcraft and Winterberg (1985) determined the stream functions and potential functions for this situation, which can be used to calculate the solute arrival time downgradient of the well. Given the
geometry shown in Figure 4, we can define the following parameters (with the well represented by a cylinder):

A well of a known radius, hydraulic conductivity and porosity is placed in a flow field of known hydraulic gradient, hydraulic conductivity, and porosity. Wheatcraft and Winterberg (1985) determined the velocity potential functions:

outside cylinder:

\[ \phi_1 = x \left( U + \frac{U' a^2}{x^2 + y^2} \right) \]  \hspace{1cm} (1)

inside cylinder:

\[ \phi_2 = U'' x \]  \hspace{1cm} (2)

where:

\[ U' = \frac{1 - k_R}{1 + k_R} \]

\[ U'' = \frac{2 U k_R}{1 + k_R} \]

x and y are spatial coordinates
where: $a =$ radius of cylinder (well)  
$k_1 =$ hydraulic conductivity outside cylinder  
$k_2 =$ hydraulic conductivity inside cylinder  
n$_1 =$ effective porosity outside cylinder  
n$_2 =$ effective porosity inside cylinder  
$k_R =$ $k_2/k_1$  
$U =$ uniform hydraulic gradient

Wheatcraft and Winterberg (1985) determined the stream functions:

outside cylinder:
$$\psi_1 = y \left[ U - \frac{U'a^2}{x^2 + y^2} \right] \quad (3)$$

inside cylinder:
$$\psi_2 = U''y \quad (4)$$

The velocity outside of the cylinder is calculated as:

$$V_{x_1} = - \frac{k_1}{n_1} \frac{\partial \phi_1}{\partial x} = \frac{\partial x}{\partial t} = - \frac{k_1}{n_1} \left[ U + \frac{U'a^2(y^2 - x^2)}{x^2 + y^2} \right] \quad (5)$$

$$V_{y_1} = - \frac{k_1}{n_1} \frac{\partial \phi_1}{\partial y} = \frac{\partial y}{\partial t} = \frac{2k_1U'a^2xy}{n_1(x^2 + y^2)^2} \quad (6)$$

Likewise, the velocity inside the cylinder is:
\[ V_{x_2} = -\frac{k_2 \phi_2}{n_2} \frac{\partial \phi_2}{\partial x} = \frac{dx}{dt} = U'' \]

Owing to \( \frac{\partial \phi_2}{\partial y} = 0 \), we know: \( V_{y_2} = 0 \)

The mass breakthrough at a downgradient control plane is calculated for numerous evenly distributed initial particle locations in the upper semi-circle; the lower semi-circle has an identical solution to the upper semi-circle. The travel time calculation is developed by first choosing increments of \( \theta \) (see Figure 5). The \( x_1 \) and \( y_1 \) locations are calculated as:

\[ x_1 = a \cos(\theta) \]  
\[ y_1 = a \sin(\theta) \]

The \( x_2 \) coordinate is the x coordinate at the downgradient control plane, while \( y_2 \) can be determined via inversion of Equation (3). The travel time from \((x_1, y_1)\) to \((x_2, y_2)\) is determined by numerical integration of Equation (6) as:

\[ t_1 = \int_{y_2}^{y_1} \frac{n_1(x^2 + y^2)^2}{2k_1 U a^2 x y} dy \]  

where \( x \) is determined from Equation (3) as:

Figure 5. Diagram of particle path.
\[ x = \left[ \frac{U'a^2y}{Uy - \psi_1} - y^2 \right]^{1/2} \]  

(10)

The travel time \( t_2 \) to \((x_1, y_1)\), at the exterior edge of the cylinder, from any lateral point in the interior of the cylinder \((x_0, y_1)\) is 

\[ t_2 = \frac{(x_1 - x_0)}{V_{x_2}} \]  

(11)

where \( V_{x_2} = U'' \). Therefore, the total travel time \( t_t \) from \((x_0, y_1)\) to \((x_2, y_2)\) is:

\[ t_t = t_1 + t_2 \]  

(12)

**HYDROCARBON RELEASE FUNCTION**

During the surface characterization of the CNTA, samples were collected from abandoned mud pits and yielded elevated levels of TPH (DOE, 1997). The highest concentrations were found in the central mud pit near the UC-1 site. The TPH consists of diesel fuel, motor oil, and crude oil, and is suspected to be from an equipment wash-rack that was adjacent to the mud pit or from mud additives, although no written documentation is available to substantiate the source. There were five mud pits used for the drilling operations at UC-4. The TPH concentrations from the UC-4 mud pits were less than those from the central mud pit and it is believed that no mud from the central mud pit was used at UC-4; however, it is possible that the same practices used at UC-1 may have been used at UC-4. Also, as previously mentioned, it was a common practice to use diesel fuel during fishing operations, such as the 30-day operation during the drilling of UC-4. Since it was not possible to retrieve in situ samples from the UC-4 borehole, samples were collected from the central mud pit for laboratory analyses.

Sampling locations were selected by targeting the highest concentrations of TPH found during the mud-pit characterization by IT Corporation (IT) (DOE, 1997). Samples were collected from two depths in four shallow hand-dug pits. Samples were stored in 1L amber glass jars at 4°C until they were analyzed. Portions of two of the samples were analyzed for physical properties. It can be seen from Table 1 that both samples have extremely similar physical characteristics. For example, the average density and average pore diameters of both samples are almost identical. In addition, both samples have similar particle size distributions and porosities. This similarity in the physical parameters of the two samples would suggest that all samples tend to have somewhat consistent and uniform physical properties. These properties are also consistent with what was known about the drilling mud. For example, it is known that the drilling mud was composed of bentonite clay. Bentonite clay is an altered volcanic ash with montmorillonite as the dominant clay mineral (Klein and Hurlbut, 1993). In general, clay minerals have high porosities and very small particle sizes, and these physical properties are manifested in both samples. In addition, montmorillonite clay has an average density of 2.5 g/cm, and it can be seen from Table 1 that both field samples have almost
identical values. Bentonite clay also has the unusual property of expanding several times its original volume when placed in water. This property explains why bentonite clays are used as a drilling mud because the clay-water mixture gives the fluid a viscosity several times that of water. It is also used for stopping leakage in soil and rocks (Klein and Hurlbut, 1993).

Table 1. A Summary of the Physical Characteristics of the Bentonite Drilling Mud Samples from the Central Mud Pit.

<table>
<thead>
<tr>
<th></th>
<th>Sample 2-2</th>
<th>Sample 4-1</th>
</tr>
</thead>
<tbody>
<tr>
<td>Average Density</td>
<td>2.50 g/cm³</td>
<td>2.49 g/cm³</td>
</tr>
<tr>
<td>Average Pore Diameter</td>
<td>162.6 µm</td>
<td>162.6 µm</td>
</tr>
<tr>
<td>Porosity</td>
<td>16.99%</td>
<td>23.37%</td>
</tr>
<tr>
<td>BET-Surface Area</td>
<td>13.2 m²/g</td>
<td>9.7 m²/g</td>
</tr>
<tr>
<td>Mean Mass Distribution</td>
<td>2.3 µm</td>
<td>5.3 µm</td>
</tr>
<tr>
<td>Mode Mass Distribution</td>
<td>2.2 µm</td>
<td>2.5 µm</td>
</tr>
<tr>
<td>Median Mass Distribution</td>
<td>0.4 µm</td>
<td>0.6 µm</td>
</tr>
</tbody>
</table>

To calculate dissolution constants of the TPH from the drilling mud, it was necessary to perform partitioning experiments as a function of time. The partitioning experiments provided data relating the equilibrium concentration of TPH in the aqueous phase to the sorbed concentration of TPH in the solid phase. To calculate the partitioning parameters, the drilling mud from a specific location was sub-sampled and divided into approximately 25 additional samples. These smaller samples (approximately 2.0 g each) were then placed into 40-mL amber glass volatile organic analysis (VOA) vials with Teflon-lined caps. Forty milliliters of synthetic groundwater was then added to the VOA vials to simulate contact with the groundwater at the CNTA. The synthetic groundwater used was made to be representative of the chemical composition of the groundwater from well HTH-1 from the Faultless site at a depth of 741 m. The 40-mL VOA vials containing the drilling mud and synthetic groundwater were then agitated on a Labquake Shaker at room temperature for increasing amounts of time.

After any one sample was sacrificed, it was then centrifuged at 3,000 rpm for 45 minutes to achieve a solid-solution separation. The aqueous phase was then removed from the solid phase and filtered using a vacuum flask and 45-micron glass-fiber filter to remove any remaining solid particles. The remaining water was then transferred into a 100-mL glass volumetric flask and filled to volume. The TPH water mixture was extracted by the addition of 3 mL of high pressure liquid chromatography-grade (HPLC-grade) hexane. After the aqueous phase had been removed from the 40-mL VOA vials, 6 mL of acetone were added to the remaining solid and the resulting mixture was then agitated for another 24 hours. The addition of 6 mL of acetone to the solid phase was intended to completely remove any remaining TPH from the solid phase. After the mixture was shaken for 24 hours, the entire contents were transferred to a 100-mL glass volumetric flask and filled to volume. The solid-water-acetone mixture was then also extracted with 3 mL of HPLC-grade hexane.

Both the equilibrium and sorbed samples were analyzed by gas chromatography (GC) to determine their TPH concentrations. The analysis method was based on U.S. Environmental...
Protection Agency (EPA) method 8015 (U.S. EPA, 1986). The TPH-hexane solutions were injected straight into the GC. The analysis was carried out by injecting 2 mL of extract using the solvent flushing technique into a GC (Hewlett Packard Series II Gas Chromatograph) equipped with a flame ionization detector. The column was a 30-m long, 0.54 cm I.D., capillary column with a film thickness of 1.2 mm (Alltech Associates, Catalog No. 955130, phase-AT-5). The oven was set to an initial temperature of 50°C and held for 1 minute. The temperature was then ramped to 260°C at the rate of 15°C/min and was held at the final temperature for 10 minutes. Inlet and detector were maintained at 270°C. An equilibration time of 1 minute between runs was allowed after the oven temperature reached 50°C. The different gases used in the GC and their flow rates are as follows: Hydrogen Flame Ionization Detector-30 ml/min; Air Flame Ionization Detector-200 ml/min; and Helium (Carrier)-2 ml/min.

The chromatograms for each sample were obtained and integrated using Hewlett Packard Chemstation Series II software. A typical example of the obtained chromatograms can be seen in Figure 6. The figure shows the GC chromatogram obtained from the equilibrium concentration of TPH after 35 days of equilibration for sample 1-1b. A standard containing straight-chain alkanes, C-9 through C-25 (Alltech Associates, Hydrocarbon kit, Cat. No. 18484) was used to calibrate the GC. A calibration curve was established for the TPH concentration in parts per million (ppm by volume) against the area under the TPH chromatogram by analyzing the standard at different lab concentrations (1,000, 100, 50, and 10 ppm). The calibration curve was then used to calculate the unknown TPH concentrations present in both the aqueous and solid phase samples.

Figure 6. GC/FID chromatogram of the 35-day equilibrium TPH concentration of sample 1-1b.
The partitioning parameters relating the concentration of TPH in the aqueous phase to the concentration of TPH in the solid phase were calculated by analyzing each sample for its TPH concentration and applying the results to the following equation:

\[ K_d = \frac{S}{C} \]  

where \( K_d \) is the distribution coefficient (m³/g); \( S \) is the sorbed concentration in grams of TPH sorbed per gram of solid (g/g); and \( C \) is the equilibrium concentration of TPH in the aqueous phase (g/m³).

The rate constants describing the dissolution of TPH from the bentonite drilling mud were also calculated using the equilibrium concentration of TPH in the aqueous phase and the original concentration of TPH in both the aqueous and solid phases. The dissolution constants were calculated by fitting the data to the following first-order equation:

\[ \frac{dC_n}{dt} = -kC \]  

where \( C_n \) is the normalized ratio of the equilibrium concentration of TPH to the original TPH concentration in both the solid and liquid phases (\( C/C_o \)); \( k \) is the first-order rate constant (day⁻¹); and \( C \) is the aqueous concentration of TPH. Integration of the previous equation yields the following linear equation:

\[ \ln \left( \frac{C}{C_o} \right) = -kt + \text{CONSTANT} \]  

where the first-order rate constant describing the dissolution of TPH from the bentonite drilling mud can then be calculated by plotting the ln \( (C/C_o) \) values versus time and then doing a best-fit linear regression of the data to determine the slope of the line.

**Laboratory Results and Discussion**

The time-dependent partitioning experiments were performed on five of the mud-pit samples from the CNTA. The samples investigated were 1-1, 1-1b, 2-2, 4-1, and 4-2. These samples were selected for study because preliminary experiments revealed that these particular mud samples contained significant quantities of TPH. In general, the time-dependent partitioning experiments revealed that the distribution coefficients for the different samples, which relate the sorbed concentration to the equilibrium concentration, decreased as time increased. For example, this type of behavior can be seen in Figure 6, which shows the distribution coefficients as a function of time for sample 1-1b. It can be seen in the figure that the distribution coefficients decrease from a maximum value of approximately 1.20 x 10⁻⁴ m³/g to a minimum value of 1.26 x 10⁻⁵ m³/g as time increases. This behavior indicates that as time increases the aqueous concentration of TPH increases for a given sample.

The distribution coefficients calculated as a function of time for each sample can be seen in the Appendix. The distribution coefficients for any one sample vary by at most two orders of
magnitude and usually by only one order of magnitude. It can also be seen from the data in the Appendix that the distribution coefficients between the various samples at similar days are also similar. This is somewhat expected given the similarity in the physicochemical characteristics of the mud-pit samples. The raw data from the chromatograms revealed, however, that the actual TPH concentrations for the samples were extremely variable. In fact, the TPH concentrations for an individual sample that had been sub-sampled were highly variable, indicating an extreme heterogeneity in the TPH concentration in the mud pit. This variability in the TPH concentration was also reported by IT in their preliminary characterization of the central mud pit (DOE, 1997).

The first-order rate constants describing the dissolution of TPH from the bentonite drilling mud can be seen in Table 2. The data that were used to calculate the rate constants for the individual samples can also be seen in the Appendix. The rate constants were calculated by a linear regression of the \( \ln \left( \frac{C}{C_0} \right) \) versus time data. It can be seen from Table 2 that the rate constants vary from a maximum value of 0.1102 per day to a minimum value of 0.0216 per day. Three of the five different samples have almost identical rate constants. This similarity in the first-order rate constants suggests that the dissolution of TPH from the bentonite drilling mud is fairly consistent even though the overall TPH concentration is extremely variable within the mud pit. This can be attributed to the similar physical characteristics of the bentonite drilling mud.

Table 2. First-Order Rate Constants for the Dissolution of TPH from Bentonite Drilling Mud.

<table>
<thead>
<tr>
<th>Sample ID</th>
<th>First-Order Rate Constant ((k)) day(^{-1})</th>
<th>Intercept</th>
<th>(R^2)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1-1</td>
<td>0.0509</td>
<td>-1.1381</td>
<td>0.64</td>
</tr>
<tr>
<td>1-1b</td>
<td>0.0216</td>
<td>-1.2079</td>
<td>0.68</td>
</tr>
<tr>
<td>2-1</td>
<td>0.0223</td>
<td>-1.2027</td>
<td>0.52</td>
</tr>
<tr>
<td>4-1</td>
<td>0.0284</td>
<td>-1.0485</td>
<td>0.38</td>
</tr>
<tr>
<td>4-2</td>
<td>0.1102</td>
<td>-1.6862</td>
<td>0.28</td>
</tr>
</tbody>
</table>

**COMPUTER MODEL DESIGN AND RESULTS**

A computer model was developed to solve the analytic solution previously described (see Equations 12-15). One-hundred Monte Carlo realizations were run from randomly chosen values of hydraulic conductivity and porosity from the alluvium and drilling-mud distributions. One-thousand particles were uniformly distributed within the mud-filled well, and the travel times for the particle to reach a control plane were monitored. An exponential release function was incorporated in the model to retard the travel times based on the TPH experimental results. Results from the model simulation are shown in Figure 7; the center breakthrough line is the mean, bounded by lines of ± two standard deviations. The first particles reached a plane 10 m from the center of the well in 262 years, the median breakthrough took 12,216 years to reach 10 m, and the final particles traveled 10 m in 33,401 years.

Flow lines are not affected greatly by the low hydraulic conductivity well at a distance of 10 m (see Figure 3); therefore, Darcian flow will control the transport of TPH from the 10 m control plane.
Figure 7. Cumulative normalized mass breakthrough of total petroleum hydrocarbon 10 m from the center of well UC-4 versus time. (Center breakthrough line is the mean, bounded by breakthrough lines of ±2 standard deviations.)

to the U.S. Department of Energy (DOE) land withdrawal boundary. Although the groundwater flow direction is not well defined in this area, in a worst-case scenario, the shortest flowpath would be essentially due south (equalling approximately 800 m). It would take an additional 45,455 years for the contaminant to travel from the 10-m plane to the DOE land boundary; however, the travel time would be much longer if advection, dispersion and sorption were taken into account. TPH travel times were affected primarily by the hydraulic conductivity and porosity of the drilling mud, then by the hydraulic properties and gradient in the alluvium, followed by the TPH release.

CONCLUSIONS

An analytical mathematic solution was developed to compute the flow velocities near the drilling-mud-filled borehole, UC-4. Contaminant transport was retarded based on the TPH release function experiments performed as part of this study. To travel 10 m, the median contaminant breakthrough occurred in about 12,000 years; the initial breakthrough took about 260 years and the final breakthrough took about 33,400 years. It would take tens of thousands of years, at a minimum, for the contaminant to travel beyond the DOE land boundary (800 m). The forces of advection, dispersion, sorption and biodegradation were not included as part of this study, but it is believed that these forces would significantly increase the travel times.
REFERENCES

Fenix and Sisson, 1968. UC-4 Hole History.


APPENDIX

TPH EXPERIMENTAL RESULTS
### Table A 1. Distribution Coefficients and ln (C/Co) Values for Sample 1-1.

<table>
<thead>
<tr>
<th>Days</th>
<th>Kd (m³/g)</th>
<th>ln (C/Co)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>1.20x10⁻⁴</td>
<td>-1.32</td>
</tr>
<tr>
<td>3</td>
<td>8.98x10⁻⁵</td>
<td>-0.687</td>
</tr>
<tr>
<td>6</td>
<td>9.66x10⁻⁵</td>
<td>-1.12</td>
</tr>
<tr>
<td>8</td>
<td>7.17x10⁻⁵</td>
<td>-0.409</td>
</tr>
<tr>
<td>10</td>
<td>7.41x10⁻⁵</td>
<td>-0.504</td>
</tr>
<tr>
<td>12</td>
<td>2.35x10⁻⁷</td>
<td>5.23</td>
</tr>
<tr>
<td>28</td>
<td>2.66x10⁻⁶</td>
<td>2.79</td>
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</table>

### Table A 2. Distribution Coefficients and ln (C/Co) Values for Sample 1-1b.

<table>
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<th>Kd (m³/g)</th>
<th>ln (C/Co)</th>
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</thead>
<tbody>
<tr>
<td>1</td>
<td>1.20x10⁻⁴</td>
<td>-1.32</td>
</tr>
<tr>
<td>3</td>
<td>8.97x10⁻⁵</td>
<td>-1.09</td>
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<tr>
<td>5</td>
<td>7.23x10⁻⁵</td>
<td>-0.997</td>
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<tr>
<td>8</td>
<td>7.17x10⁻⁵</td>
<td>-0.919</td>
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<tr>
<td>10</td>
<td>7.41x10⁻⁵</td>
<td>-0.977</td>
</tr>
<tr>
<td>20</td>
<td>4.02x10⁻⁵</td>
<td>-0.634</td>
</tr>
<tr>
<td>22</td>
<td>9.52x10⁻⁵</td>
<td>-1.09</td>
</tr>
<tr>
<td>24</td>
<td>6.38x10⁻⁵</td>
<td>-0.786</td>
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<tr>
<td>35</td>
<td>1.26x10⁻⁵</td>
<td>-0.252</td>
</tr>
<tr>
<td>44</td>
<td>1.57x10⁻⁵</td>
<td>-0.280</td>
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### Table A 3. Distribution Coefficients and ln (C/Co) Values for Sample 2-2.

<table>
<thead>
<tr>
<th>Days</th>
<th>Kd (m³/g)</th>
<th>ln (C/Co)</th>
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</thead>
<tbody>
<tr>
<td>1</td>
<td>8.76x10⁻⁵</td>
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<td>3</td>
<td>1.71x10⁻⁴</td>
<td>-1.73</td>
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<td>8</td>
<td>2.05x10⁻⁵</td>
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<td>33</td>
<td>5.46x10⁻⁵</td>
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</tr>
<tr>
<td>35</td>
<td>1.77x10⁻⁵</td>
<td>-0.326</td>
</tr>
<tr>
<td>37</td>
<td>2.14x10⁻⁵</td>
<td>-0.379</td>
</tr>
<tr>
<td>40</td>
<td>1.03x10⁻⁵</td>
<td>-0.222</td>
</tr>
</tbody>
</table>
Table A 4. Distribution Coefficients and ln (C/Co) Values for Sample 4-1.

<table>
<thead>
<tr>
<th>Days</th>
<th>Kd (m$^3$/g)</th>
<th>ln (C/Co)</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>1.54x10^{-4}</td>
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<td>3</td>
<td>5.78x10^{-5}</td>
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<tr>
<td>7</td>
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<td>10</td>
<td>6.75x10^{-5}</td>
<td>-0.863</td>
</tr>
<tr>
<td>22</td>
<td>1.91x10^{-5}</td>
<td>-0.400</td>
</tr>
<tr>
<td>24</td>
<td>3.77x10^{-5}</td>
<td>-0.511</td>
</tr>
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Table A 5. Distribution Coefficients and ln (C/Co) Values for Sample 4-2.

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<th>Days</th>
<th>Kd (m$^3$/g)</th>
<th>ln (C/Co)</th>
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</thead>
<tbody>
<tr>
<td>1</td>
<td>1.113x10^{-4}</td>
<td>-1.02</td>
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<td>2</td>
<td>4.81x10^{-4}</td>
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<td>7.47x10^{-5}</td>
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<td>8</td>
<td>1.53x10^{-4}</td>
<td>-1.23</td>
</tr>
<tr>
<td>10</td>
<td>1.82x10^{-5}</td>
<td>-0.268</td>
</tr>
</tbody>
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DISTRIBUTION

Raymon D. Cox, Acting, Director
Contract Management Division
Nevada Operations Office
U.S. Department of Energy
P.O. Box 98518
Las Vegas, NV 89193–8518

Leslie A. Monroe
Environment, Safety & Health Division
Nevada Operations Office
U.S. Department of Energy
P.O. Box 98518
Las Vegas, NV 89193–8518

Beverly Colbert
Contract Management Division
Nevada Operations Office
U.S. Department of Energy
P.O. Box 98518
Las Vegas, NV 89193–8518

Mike Sabbe
Bechtel Nevada
P.O. Box 98521
Las Vegas, NV 89193-8521

Doug Duncan
Hydrology Program Manager
Environment, Safety & Health Division
Nevada Operations Office
U.S. Department of Energy
P.O. Box 98518
Las Vegas, NV 89193–8518

Monica Sanchez
Environmental Restoration Division
Nevada Operations Office
U.S. Department of Energy
P.O. Box 98518
Las Vegas, NV 89193–8518

Kenneth A. Hoar, Director
Environment, Safety & Health Division
Nevada Operations Office
U.S. Department of Energy
P.O. Box 98518
Las Vegas, NV 89193–8518

Peter Sanders
Environmental Restoration Division
Nevada Operations Office
U.S. Department of Energy
P.O. Box 98518
Las Vegas, NV 89193–8518

Marjory Jones
Desert Research Institute
Water Resources Center
2215 Raggio Parkway
Reno, NV 89512-1095

David Shafer
Water Resources Center
Desert Research Institute
755 E. Flamingo Road
Las Vegas, NV 89119-7363

David K. Smith
Isotopes Sciences Division
Lawrence Livermore National Laboratory
P.O. Box 808, M/S L231
Livermore, CA 94550

Randy Laczniak
U.S. Geological Survey
Water Resources Division
6770 S. Paradise Rd.
Las Vegas, NV 89119

Doug Trudeau
U.S. Geological Survey
Water Resources Division
6770 S. Paradise Rd.
Las Vegas, NV 89119

Tim White
International Technology Corporation
P.O. Box 93838, M/S 439
Las Vegas, NV 89193–3838